

Continuous streamflow simulation for index flood estimation in an Alpine basin of Northern Italy



ENVIRON

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Aim of the work



- Flood frequency estimation is crucial in both engineering practice and hydrologic research
- In ungauged or poorly gauged river basins direct methods can not be used
- We present a reconstruction of continuous streamflow using physically based spatially distributed hydrological model for index flood estimation



The index flood method

For
homogeneous
region

$$q_T = x_T \cdot q_{index}$$

q_T
T-year quantile
of flood flows

x_T
T-year quantile of
normalized flood
flows in the region,
growth factor

q_{index}
index flood

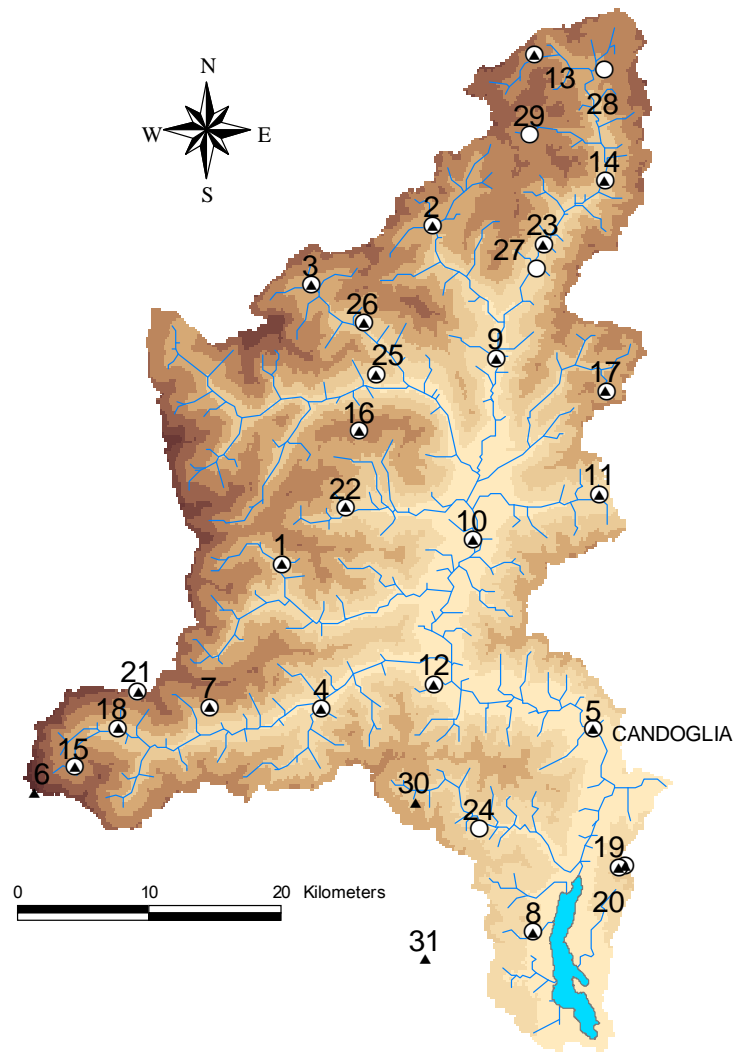


generally assumed
equal to the mean
maximum annual
flood

Bocchiola, D., De Michele, C., and Rosso, R., (2003).
"Review of recent advances in index flood
estimation." *Hydrol. Earth Syst. Sc.*, 7(3), 283-296.



The case study: Toce river



- ▲ Thermometers
- Rain gauges
- DEM (m a.s.l.)
- 194 - 679
- 680 - 1165
- 1166 - 1651
- 1652 - 2137
- 2138 - 2623
- 2624 - 3109
- 3110 - 3595
- 3596 - 4081
- 4082 - 4567

Area: 1800 km²

Alpine basin: nearly 32% of the total area above 2000 m a.s.l..

28 rain gauges

27 thermometers

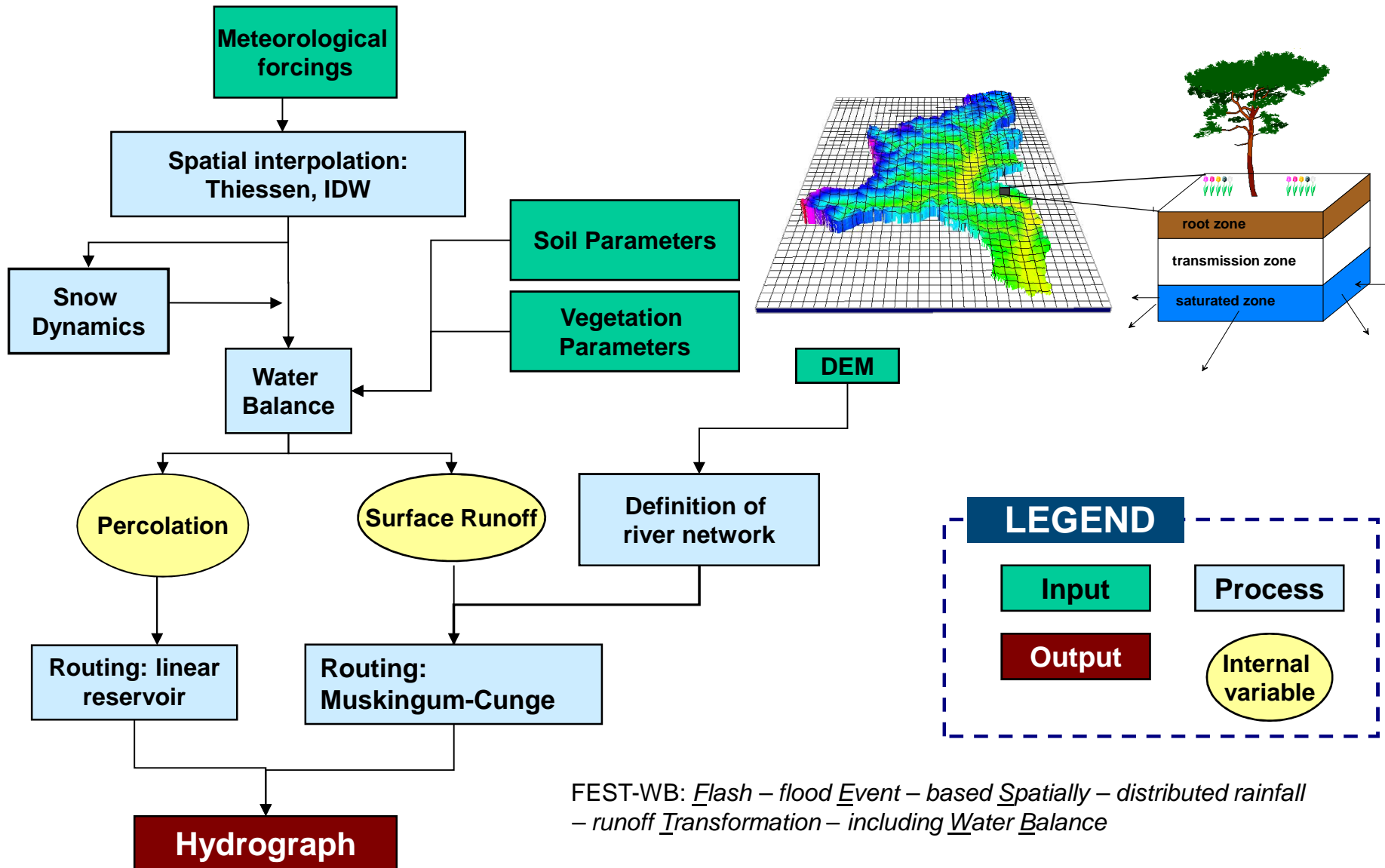
Number of available station is variable from 1982

At Candoglia (1500 km²) 53 years of max annual flood (1933-1998) and hourly discharge measurements from 2000-2010

$$Q_{index} = 1076 \text{ m}^3/\text{s}$$

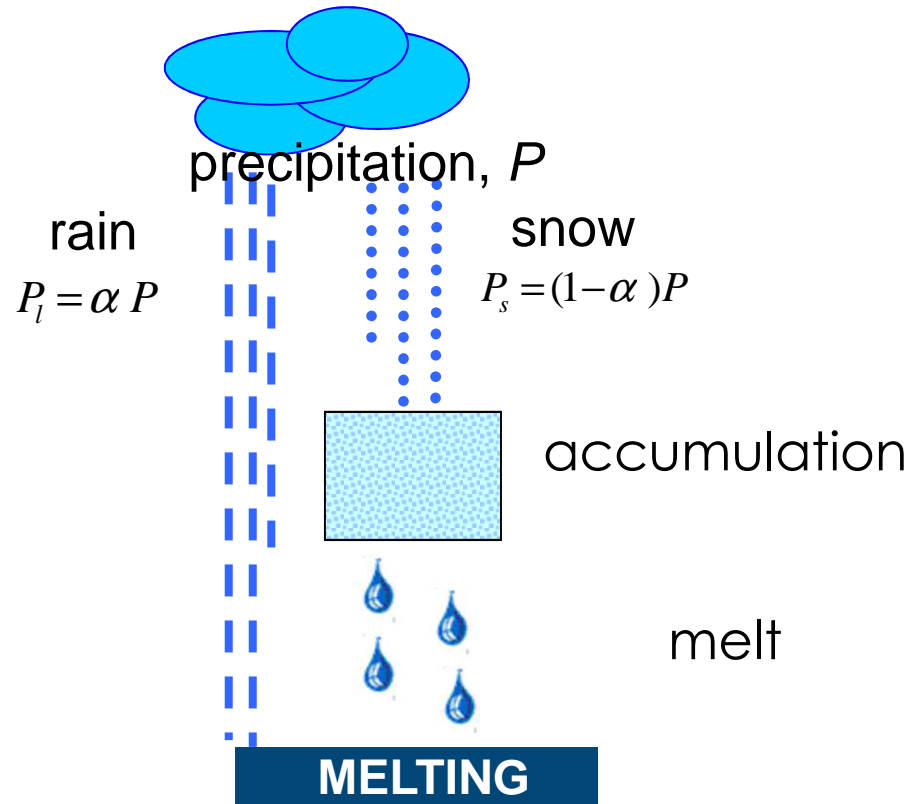


Hydrological model FEST-WB





Snow dynamics



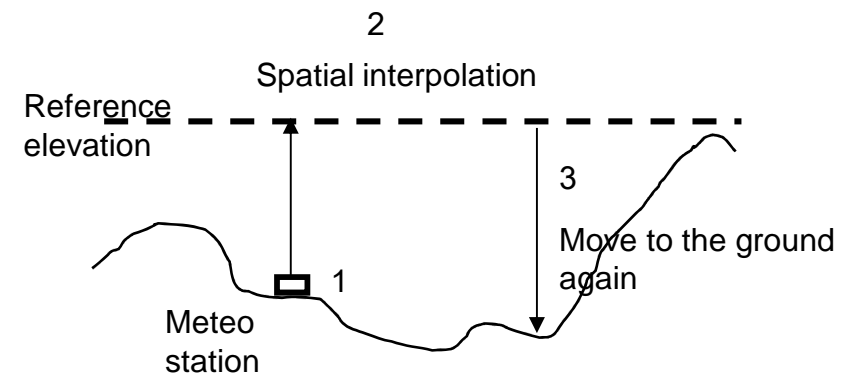
$$\text{SnowMelt} = C_m (T_a - T_b)$$

T_a = air temperature

T_b = threshold temperature

C_m = melt coefficient [m/(s °C)]

Air temperature spatial interpolation



- (1) Move measurements on a reference plane keeping into account a fixed thermal gradient ($-0.0065 \text{ °C m}^{-1}$)
- (2) Spatial interpolation on the reference plane
- (3) Data are taken to the ground keeping into account a fixed thermal gradient ($-0.0065 \text{ °C m}^{-1}$).



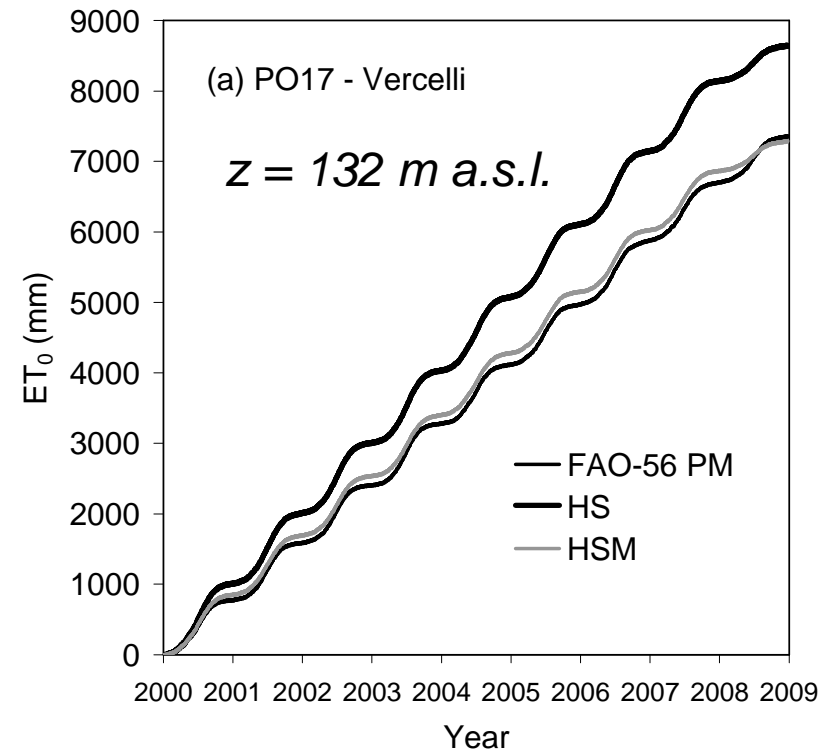
Hydrological model FEST-WB

Modified Hargreaves equation
keeping into account elevation

Evapotranspiration

$$ET_{0,HSM} = (0.817 + 0.00022 \cdot z) \cdot HC \cdot R_a \cdot (T_{max} - T_{min})^{HE} \left(\frac{T_{max} + T_{min}}{2} + HT \right)$$

	HS	HS elev
upper Po		
MBE	0.154	-0.030
RMSE	0.885	0.784
Rhone		
MBE	-0.145	0.031
RMSE	0.670	0.655
upper Po + Rhone		
MBE	0.079	-0.014
RMSE	0.836	0.754



Ravazzani, G., Corbari, C., Morella, S., Gianoli, P., Mancini, M. (2011), Modified Hargreaves-Samani equation for the assessment of reference evapotranspiration in Alpine river basins. *J. Irrig. Drain. E-ASCE*, accepted, 10.1061/(ASCE)IR.1943-4774.0000453 available online

Reduces overestimation at low elevation

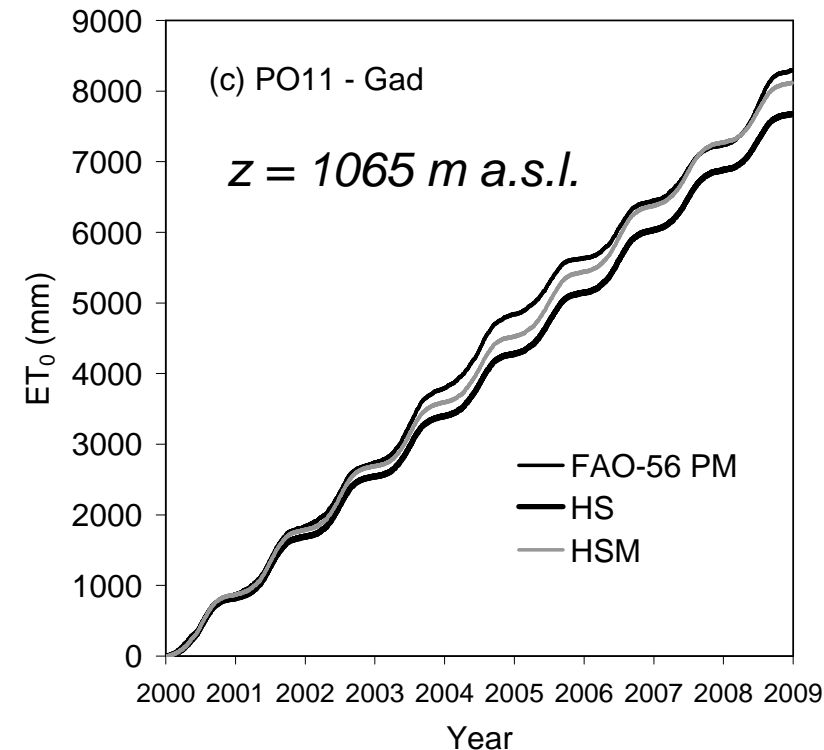


Modified Hargreaves equation
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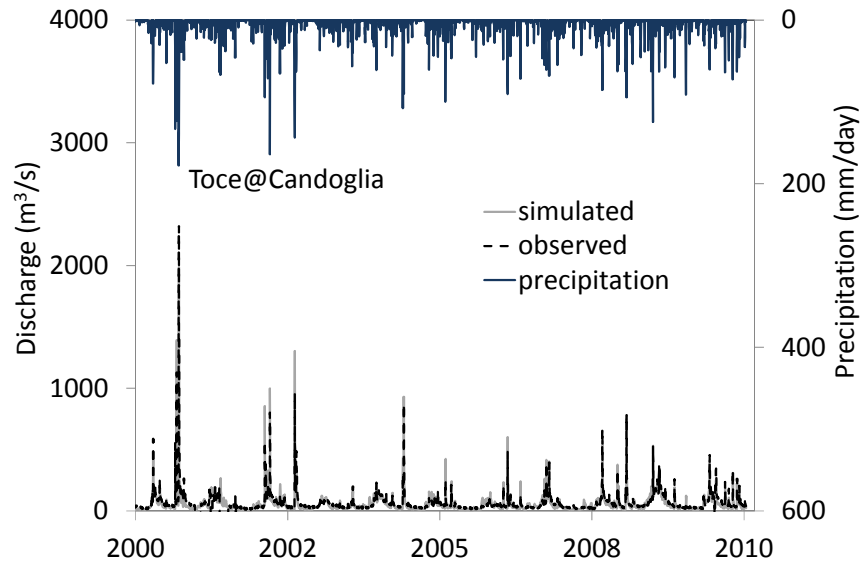


Ravazzani, G., Corbari, C., Morella, S., Gianoli, P., Mancini, M. (2011), Modified Hargreaves-Samani equation for the assessment of reference evapotranspiration in Alpine river basins. *J. Irrig. Drain. E-ASCE*, accepted, 10.1061/(ASCE)IR.1943-4774.0000453 available online

Reduces underestimation at high elevation



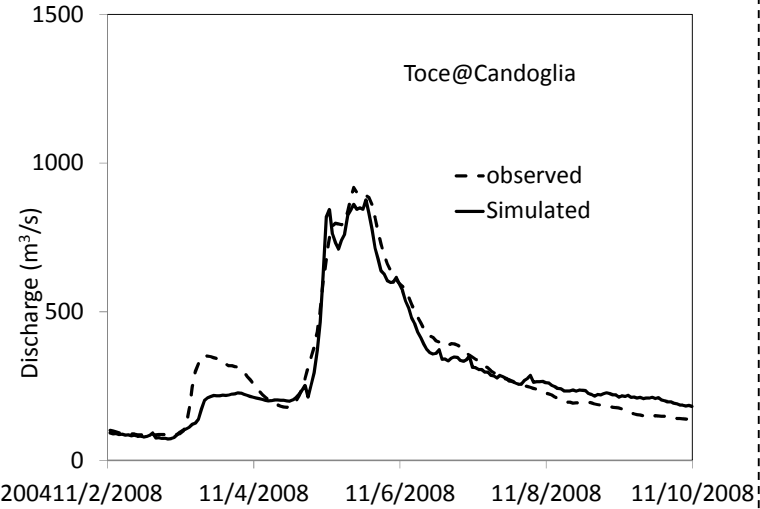
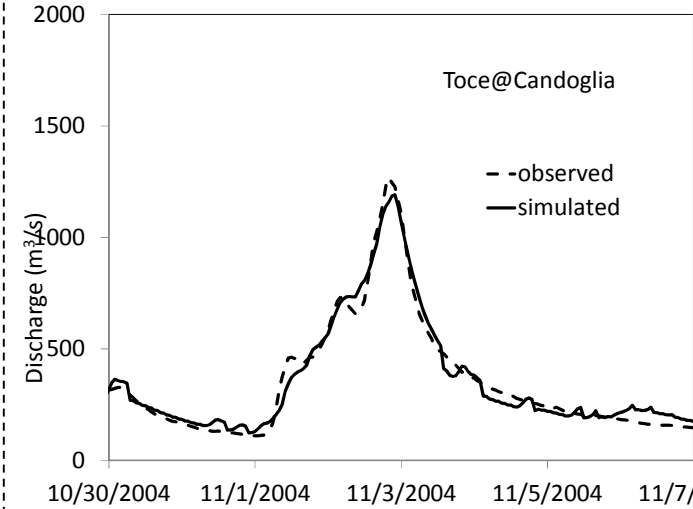
Model validation



DAILY

Period	RMSE	Nash-Sutcliffe
2000-2010	29.53	0.88

HOURLY





max annual flood

Simulated years: 29

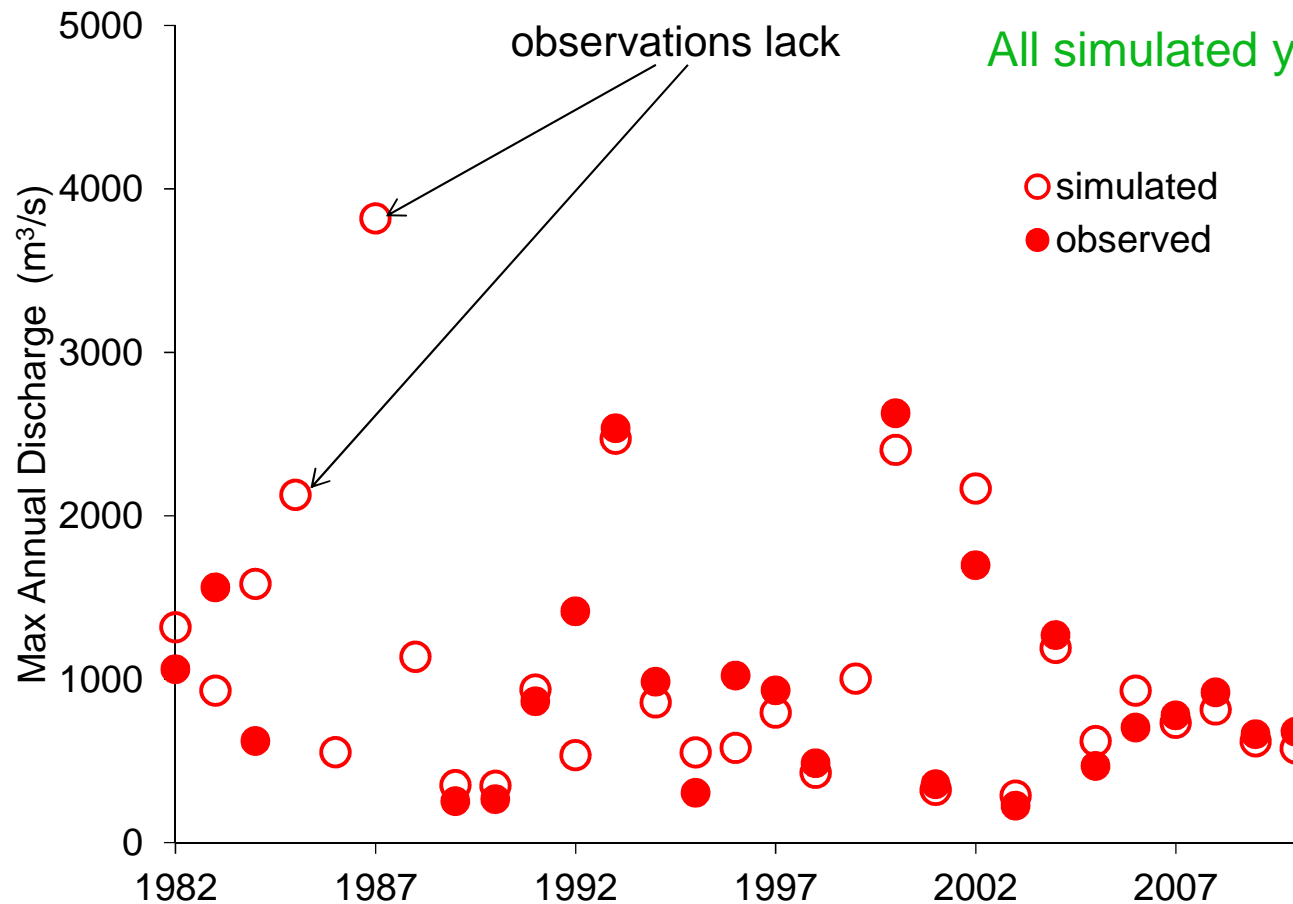
Available observations: 24

INDEX FLOOD ESTIMATION

All observations: 945 m³/s

24 Selected simulated years: 935 m³/s

All simulated years: 1070 m³/s





- The FEST-WB distributed hydrological model proved to be accurate in simulating flood events on river Toce even without calibration
- Index flood estimation based on continuous streamflow simulation is useful in ungauged or poorly gauged basin
- Index flood estimation based on continuous streamflow simulation is useful for reconstruction of missing data so to reduce probability of underestimating index flood



Acknowledgments



POLITECNICO
DI MILANO

ENVIRON

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THANK YOU FOR YOUR ATTENTION !

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